

Tropospheric delay in microwave propagation in Nigeria.

Samuel T. Ogunjo*

Department of Physics, Federal University of Technology, Akure, Ondo State, Nigeria

Joseph B. Dada

Department of Physics, Federal University of Technology, Akure, Ondo State, Nigeria

Sunday S. Oluyamo

Department of Physics, Federal University of Technology, Akure, Ondo State, Nigeria

Ibiyinka A Fuwape

Department of Physics, Federal University of Technology, Akure, Ondo State, Nigeria

Abstract

Satellite communication systems suffer from the systematic error of tropospheric delay. Accurate estimation of this delay is essential for communication budget and planning. This study investigates the tropospheric delay in three Nigeria cities: Abuja, Lagos, Port-Harcourt using two different models (Saastominen and Hopfield). Three year atmospheric data for surface pressure, relative humidity and temperature obtained at 5-mins interval were acquired from the Tropospheric Data Acquisition Network (TRODAN) archives. Computed radio refractivity values showed distinct seasonal de-

*Corresponding author

Email addresses: stogunjo@futa.edu.ng (Samuel T. Ogunjo),
babatunde.dada@elizadeuniversity.edu.ng (Joseph B. Dada),
ssoluyamo@futa.edu.ng (Sunday S. Oluyamo), iafuwape@futa.edu.ng (Ibiyinka A Fuwape)

¹Elizade University, Ilara-Mokin, Ondo State, Nigeria

²Michael and Cecilia Ibru University, Ughelli, Delta State, Nigeria

pendence in Abuja with low and high values during the dry and wet season respectively. The Hopfield model predicts higher hydrostatic delay values than the Saastominen model. In the non-hydrostatic delay, the two models converge to a single values at high temperature. Theorems were proposed with proofs to explain the relationship observed between the two models.

Keywords: TRODAN, Saastominen, Hopfield, tropospheric delay, radio refractivity

1. Introduction

Satellite communication systems, such as the Global Positioning System, involves transmission of information using electromagnetic waves from a satellite to ground based stations through the atmosphere ([Abdelfatah et al., 2015](#)). The different layers of the troposphere causes delay and refraction of the signal as it passes through. The microwave propagation through the atmosphere experience tropospheric delay due to electrically neutral of the atmosphere and it is completely independent of the signal frequency ([Opaluwa et al., 2013](#); [Younes, 2016](#)). The main meteorological parameters that affect propagation through the troposphere includes relative humidity, atmospheric temperature, and atmospheric pressure. The path traveled through by a propagated signal in the atmosphere contributed to the tropospheric propagation delays. Tropospheric delay is the variability effect of refractive index on radio signal traveling through the electrically-neutral atmosphere. This propagation delay is commonly determine by small changes in refractive index describe by the term called refractivity (N) ([Adegoke and Onasanya,](#)

2008)

$$N = (n - 1) \times 10^6 \quad (1)$$

where n is the an atmospheric varying index.

According to [Mendes and Langley \(1995\)](#), mismodeling of tropospheric delay in radio wave propagation is a significant source of error in space geodesy techniques. This delay is computed as the difference between the path taken by the signal and the path the signal would have taken in a vacuum. Tropospheric delay has two components - Hydrostatic (dry) delay and non-hydrostatic (wet) delay. The dry component of tropospheric delay accounts for 80 - 90% of the total delay. Wet delay is used in obtaining precipitable water vapour. The non-hydrostatic component, although a small component of the total delay, is difficult to estimate due to the high variability of atmospheric water vapour ([Liu et al., 2017](#)).

In tropical Africa, refractivity of signals has been studied extensively. However, there is a dearth of literature and research on tropospheric delay in the region. This study investigates the tropospheric delay at three cities in Nigeria using two models: Saastamoinen ([Saastamoinen, 1972](#)) and Hopfield ([Hopfield, 1969](#)). Modelling of tropospheric delay will help mitigate the influence of the atmosphere on communication systems.

2. Methodology

Data for this study was obtained from the Tropospheric Data Acquisition Network (TRODAN) which monitor, collect and provide real time meteorological data of the lower atmosphere which covers region from the surface of the Earth to the altitude of about 11 km from different locations across Nige-

ria using the Campbell Scientific Automatic Weather Station. Daily data for temperature, relative humidity and surface pressure for three years (April 1, 2008 - March 31, 2011) recorded at 5 minutes interval were retrieved from the archives of TRODAN. Statistics for the three locations considered in this study are presented in Table 1.

Table 1: Geographical statistics of study locations.

Location	Latitude ($^{\circ}N$)	Longitude ($^{\circ}E$)	Altitude (m)
Abuja	9.0667	7.4833	536
Lagos	6.4343	3.3226	7
Port Harcourt	4.7848	6.9918	20

Temperature (T) in Kelvins and Surface pressure (P) in hPa were used to compute the Hopfield (Hopfield, 1969) and Saastamoinen (Saastamoinen, 1972) models. The Hopfield Dry delay model is given by

$$Z_D^H = \frac{0.62291}{T^2} + 0.0023081P \quad (2)$$

Hopfield wet delay model is given by

$$Z_W^H = 0.07402 \cdot \frac{e}{T^2} \cdot H_T \quad (3)$$

where H_T is the height of the tropopause which is taken to be 12km.

Saastamoinen hydrostatic delay is computed

$$Z_D^S = \frac{0.0022767P}{1 - 0.00266 \cos 2\phi - 0.00028h} \quad (4)$$

where ϕ and h are the ellipsoidal latitude and surface height above the ellipsoid in km, respectively. And the wet component given by

$$Z_W^S = 0.002277 \left(\frac{1255}{T} + 0.05 \right) e \quad (5)$$

e is the partial water vapour pressure at the surface computed from the expression

$$e = \frac{5854}{T^5} 10^{20 - \frac{2950}{T}} \quad (6)$$

Total refractivity is computed from the expression

$$N = 77.6 \frac{P}{T} + 6.48 \frac{e}{T} + 3.75 \times 10^5 \frac{e}{T} \quad (7)$$

3. Results and Discussion

The temporal variation of dry and wet component of refractivity as well as the total refractivity over the study locations is presented in Figure 1. The dry component ranges from 245 - 270 N-units and constitute the larger part of the total refractivity. Lagos and Port Harcourt showed similar values and trend in the dry component of refractivity. This can be attributed to the two locations being coastal regions. Although not pronounced, Lagos and Port Harcourt also showed seasonal variations as higher values of refractivity were observed during the wet season and low values in the dry season. Abuja, being an inland station, exhibit lower values than the coastal cities of Lagos and Port Harcourt. This is attributed to the high value of water vapour in the atmosphere of the coastal cities. The computed wet component of radio refractivity showed remarkable features. During the wet season, the three stations were found to have similar values which is between

110 and 130. However, during the dry season values of wet refractivity for Abuja dropped. The wet component of radio refractivity was found to be a significant component of the total refractivity (Fuwape et al., 2016).

Figure 2 shows the temporal variation of the hydrostatic zenith delay for the three locations under consideration using the Saastomoinen and Hopfield models. For the three locations considered, the Hopfield model suggests an higher value for hydrostatic delay than the Saastomoinen model, albeit, the two models were found to track each other. The difference between Hopfield and Saastomoinen hydrostatic model can be accounted for by considering Theorem 1. The difference between the models become smallest at high temperature. Abuja showed the lowest hydrostatic delay compared to the coastal cities of Lagos and Port Harcourt. This result confirms high hydrostatic delay along coastal regions and low values in inland stations as reported by (Fuwape et al., 2016).

Theorem 1. *The difference between the Saastomoinen and Hopfield hydrostatic delay can be expressed as $\frac{0.62291}{T^2} + 3.14 \times 10^{-5}P$*

Proof. Subtracting the Hopfield hydrostatic model (Equation 2) from the Saastomoinen hydrostatic model (Equation 4), we obtain

$$Z_D^H - Z_D^S \approx \frac{0.62291}{T^2} + 3.14 \times 10^{-5}P \quad (8)$$

At low altitude ($h < 1km$), $0.00028h \approx 0$. Also, $\max(0.00266 \cos 2\phi = 0.00266)$. Hence, we assumed $1 - 0.00266 \cos 2\phi - 0.00028h \approx 1$. \square

Non-hydrostatic delay (Figure 3) showed seasonal variation. Values of the non-hydrostatic delay were found to be low during the wet season and

high during the dry season. Values of non-hydrostatic delay were found to be closely related but not identical for Lagos and Port-Harcourt. However, identical values were found for the two models in Abuja during the dry season in the temporal variation of non-hydrostatic delay. The diurnal variation of non-hydrostatic and hydrostatic delay are shown in Figures 4 and 5 respectively. The values of the dry and wet delays were found to rise from the low values in the morning to maximum at mid-day. The hydrostatic and non-hydrostatic delay for Abuja was found to peak at 1500 local time while Lagos and Port-Harcourt peaked at 1400 hours local time. In Abuja, the two models converged to the same values as it approaches mid-day in the non-hydrostatic delay. The smallest differences between the two models in the hydrostatic delay were observed in Abuja. To account for the convergence between the Saastominen and Hopfield wet delay models during the dry season and mid-day, we propose Theorem 2. During the dry season and around mid-day, the temperature in Abuja can rise beyond 307 K, hence, the convergence of the two models.

Theorem 2. *Given $H_T = 12km$, the Saastominen and Hopfield non-hydrostatic delay are equivalent when temperature equals 307K*

Proof. Equating the non-hydrostatic Saastominen (Equation 5) and Hopfield (Equation 3), we obtain

$$2.857635T + 0.00011385T^2 = 0.07402H_T \quad (9)$$

Assuming $H_T = 12km$ and solving the quadratic equation yields $T = 307K$ □

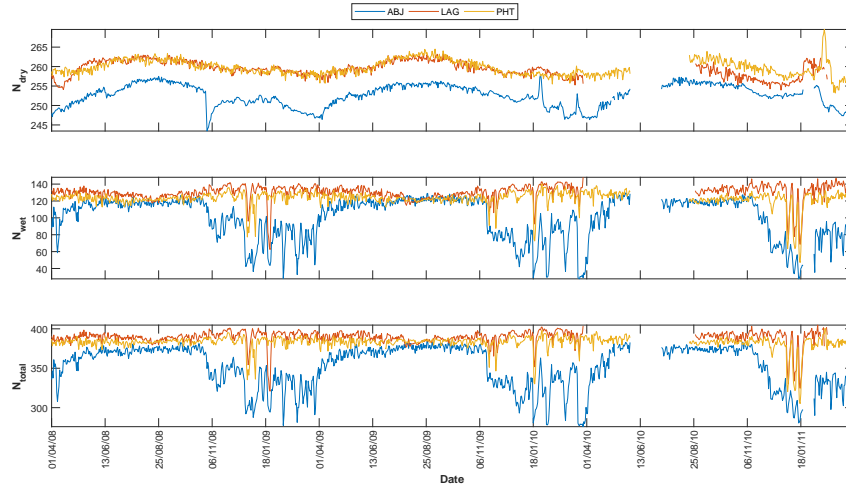


Figure 1: Temporal variation of dry component of radio refractivity (top panel), wet refractivity (middle panel) and total radio refractivity (bottom panel).

4. Conclusion

In this paper, radio refractivity and tropospheric delay for three tropical locations within Nigeria have been investigated. It was established that the coastal climate of Lagos and Port-Harcourt is responsible for the similar trend and values in the dry, wet and total refractivity. The refractivity values for Abuja, an inland city, was found to have lower values than that of the coastal cities. This factor is also responsible for the visible seasonal behaviour of refractivity in Abuja. The temporal variation of hydrostatic and non-hydrostatic tropospheric delay was also studied in this research using the Saastominen and Hopfield models. The difference between both models for the hydrostatic delay was attributed to the temperature dependence of both models. In the non-hydrostatic delay, the two models were found to have similar values at a temperature of $33^{\circ}C$. Proofs were presented to established

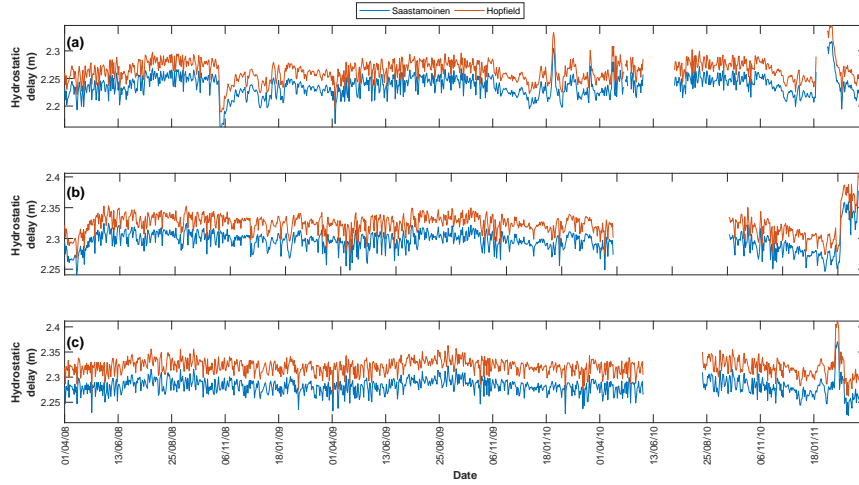


Figure 2: Temporal variation of hydrostatic zenith delay for (a) Abuja (b) Lagos (c) Port Harcourt.

the proposed theorems for hydrostatic and non-hydrostatic delays.

Acknowledgements

The results presented in this paper rely on TRODAN data collected and managed by the Centre for Atmospheric Research, National Space Research and Development Agency, Federal Ministry of Science and Technology, Anyigba, Nigeria. We thank the Centre for Atmospheric Research and their partners for promoting high standards of atmospheric observatory practice as well as the Federal Government of Nigeria for continuous funding of the Nigerian Space programme (www.carnasrda.com).

References

M. Abdelfatah, A. E. Mousa, G. S. El-Fiky, Precise troposphere delay model for Egypt, as derived from radiosonde data, NRIAG Journal of Astronomy

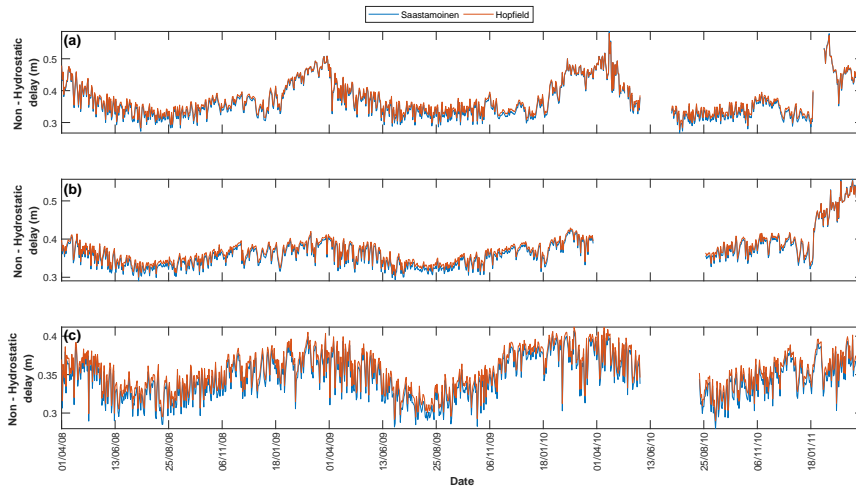


Figure 3: Temporal variation of non-hydrostatic delay for (a) Abuja (b) Lagos (c) Port Harcourt.

and Geophysics 4 (1) (2015) 16–24.

Y. D. Opaluwa, Q. A. Adejare, Z. A. T. Suleyman, I. C. Abazu, T. O. Adewale, A. O. Odesanmi, V. C. Okorochoa, Comparative analysis of five standard dry tropospheric delay models for estimation of dry tropospheric delay in GNSS positioning, *American Journal of Geographic Information System* 2 (4) (2013) 121–131.

S. A.-M. Younes, Modeling investigation of wet tropospheric delay error and precipitable water vapor content in Egypt, *The Egyptian Journal of Remote Sensing and Space Science* 19 (2) (2016) 333–342.

A. Adegoke, M. Onasanya, Effect of propagation delay on signal transmission, *Pacific J Sci Technol* 9 (2008) 13–19.

V. Mendes, R. Langley, Zenith wet tropospheric delay determination using

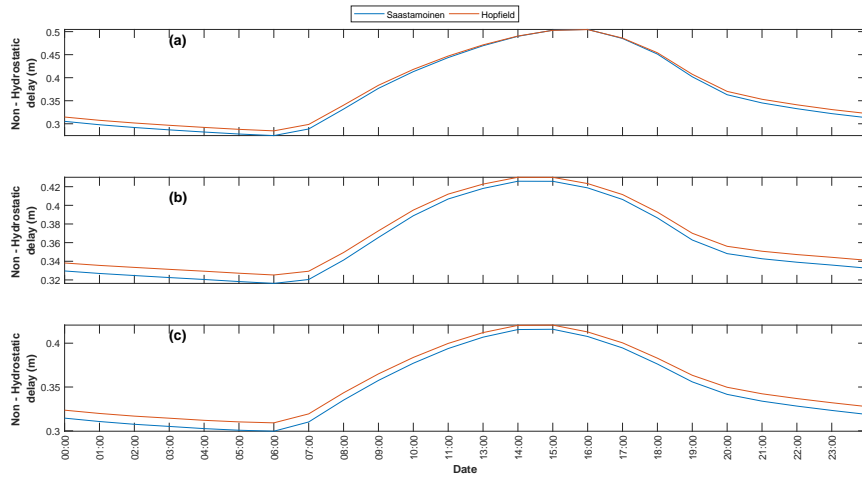


Figure 4: Diurnal variation of non-hydrostatic delay for (a) Abuja (b) Lagos (c) Port Harcourt.

prediction models: Accuracy analysis, *Cartografia e Cadastro* 2 (1995) 41–47.

J. Liu, X. Chen, J. Sun, Q. Liu, An analysis of GPT2/GPT2w+ Saastamoinen models for estimating zenith tropospheric delay over Asian area, *Advances in Space Research* 59 (3) (2017) 824–832.

J. Saastamoinen, Atmospheric correction for the troposphere and stratosphere in radio ranging satellites, *The use of artificial satellites for geodesy* 15 (1972) 247–251.

H. Hopfield, Two-quartic tropospheric refractivity profile for correcting satellite data, *Journal of Geophysical research* 74 (18) (1969) 4487–4499.

I. A. Fuwape, S. T. Ogunjo, J. B. Dada, G. A. Ashidi, I. Emmanuel, Phase synchronization between tropospheric radio refractivity and rainfall

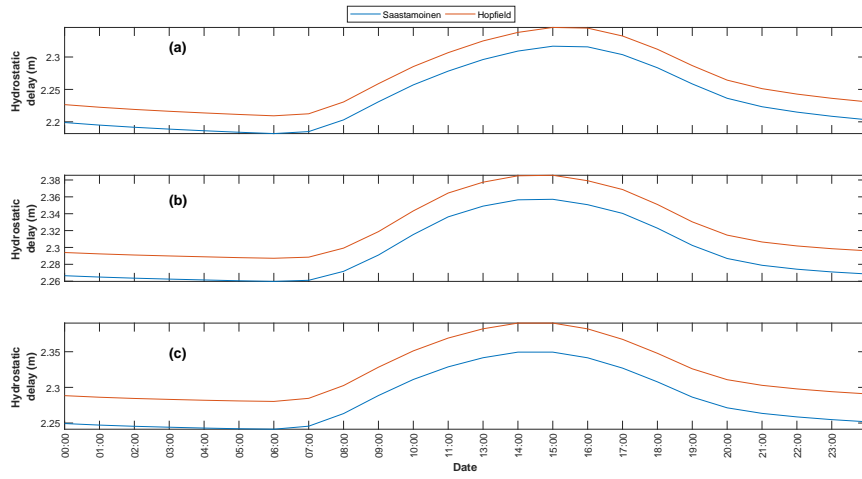


Figure 5: Diurnal variation of hydrostatic delay for (a) Abuja (b) Lagos (c) Port Harcourt.

amount in a tropical region, *Journal of Atmospheric and Solar-Terrestrial Physics* 149 (2016) 46–51.